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1 The fate of polar trace organic compounds in the hyporheic zone

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12 **Abstract**

13 The hyporheic zone (HZ) is often considered to efficiently remove polar trace organic compounds
14 (TrOCs) from lotic systems, mitigating potential adverse effects of TrOCs on ecosystem functioning and
15 drinking water production. Predicting the fate of TrOCs in the hyporheic zone (HZ) is difficult as the in-
16 situ removal rate constants are not known and the biogeochemical factors as well as hydrological
17 conditions controlling the removal efficiency are not fully understood. To determine the in-situ removal
18 efficiency of the HZ for a variety of TrOCs as a function of the biogeochemical milieu, we conducted a
19 field study in an urban river near Berlin, Germany. Subsurface flow was studied by time series of
20 temperature depth profiles and the biogeochemical milieu of the HZ by concentration depth profiles.
21 These results, in conjunction with a 1D advection-dispersion transport model, were used to calculate first-
22 order removal rate constants of several polar TrOCs in the HZ. For the majority of TrOCs investigated,
23 removal rate constants were strongly dependent on redox conditions, with significantly higher removal
24 rates observed under predominantly suboxic (i.e. denitrifying) compared to anoxic (i.e. Fe and Mn
25 reducing) conditions. Compared to previous studies on the fate of TrOCs in saturated sediments, half-lives
26 within oxic/suboxic sections of the HZ were relatively low, attributable to the site-specific characteristics
27 of the HZ in a stream dominated by wastewater treatment plant effluent. For nine out of thirteen
28 investigated TrOCs, concentrations decreased significantly in the HZ with relative removal percentages
29 ranging from 32% for primidone to 77% for gabapentin. For many TrOCs, removal efficiency decreased
30 drastically as redox conditions became anoxic. For the majority of compounds investigated here, the HZ
31 indeed acts as an efficient bioreactor that is capable of removing TrOCs along relatively short flow paths.
32 Depending on the TrOC, removal capacity may be enhanced by either increasing the magnitude of
33 groundwater-surface exchange fluxes, by increasing the total residence time in the HZ or the exposure
34 time to suboxic zones, respectively.

35 **Keywords**

36 Pharmaceuticals, organic micropollutants, biotransformation, urban water cycle, removal efficiency

37 **Introduction**

38 Trace organic compounds (TrOCs) derived from wastewater treatment plants (WWTP) are frequently
39 detected in surface waters in which they impair ecosystem functioning (Schwarzenbach et al., 2006) and
40 pose risks for drinking water production, especially in urban areas where water cycles might be partially
41 closed (Pal et al., 2014). The hyporheic zone (HZ), the interface between groundwater and surface water
42 in streams, has often been considered to efficiently remove TrOCs (Lewandowski et al., 2011). Depending
43 on the magnitude and direction of groundwater-surface water exchange and hyporheic exchange fluxes,
44 the HZ may contribute to the overall attenuation of TrOCs in streams (Riml et al., 2013; Writer et al.,
45 2013) and may act as a reactive barrier protecting groundwater from surface water pollution and vice
46 versa (Huntscha et al., 2013).

47 Quantitative information on the reactivity of TrOCs in saturated sediments, often expressed as first-order
48 removal rate constants, predominantly originates from either laboratory experiments (Burke et al., 2014;
49 Li et al., 2015; Radke and Maier, 2014) or bank filtration studies (Henzler et al., 2014; Huntscha et al.,
50 2013). For many TrOCs, reported removal rate constants were found to be highly site- and experiment
51 specific and varied up to three orders of magnitude among different subsurface environments and
52 experimental setups (Greskowiak et al., 2017). It has further been shown that a variety of physical and
53 biogeochemical factors, such as temperature (Burke et al., 2014), redox conditions (Wiese et al., 2011),
54 concentrations of labile dissolved organic carbon (DOC) (Hoppe-Jones et al., 2012) and the abundance
55 and diversity of microorganisms (Bertelkamp et al., 2016) may influence the reactivity of TrOCs in
56 saturated sediments.

57 Compared to other subsurface environments, hyporheic zones are typically characterized by intense
58 microbial activity, diverse microbial communities and steep redox gradients (Boano et al., 2014). Field
59 investigations that link the fate of TrOCs to the distinct biogeochemical characteristics in the HZ, are
60 widely lacking and thus, in-situ removal rate constants of TrOCs in the HZ are not known. Removal
61 efficiency of the HZ is generally not only a function of the removal rate constants and thus the

62 biogeochemical milieu, but also of the time over which an attenuation reaction can take place (i.e. the
63 residence time). The spatial distribution of biogeochemical parameters, such as the ambient redox
64 potential in the HZ, is often tightly coupled to flow characteristics. Compared to bank filtration systems,
65 residence times in the HZ are usually much shorter. The in-situ interactions between transport timescales
66 in the HZ and the biogeochemical conditions that control the ambient reactivity of TrOCs in HZ, (i.e. the
67 factors that govern the efficiency of the HZ in removing polar TrOCs) are poorly understood.

68 The present study had two aims: (i) providing in-situ first-order removal rate constants for several TrOCs
69 for the HZ of an urban lowland stream and (ii) determine the hydrological and biogeochemical factors that
70 limit the efficiency of the HZ in removing TrOCs. To this end, in-situ first-order removal rate constants
71 are determined for a variety of polar WWTP derived TrOCs under different redox conditions.

72 We hypothesize that in a river that is dominated by WWTP effluent removal rate constants in the
73 hyporheic zone substantially deviate from rates obtained in laboratory studies and bank filtration systems.
74 As reaction rates for many compounds are likely to be redox-dependent, reaction rates are expected to be
75 strongly controlled by redox conditions in the HZ. Consequently, removal efficiency is likely to depend on
76 both the extent of redox zones as well as on the magnitude of groundwater-surface exchange and
77 hyporheic exchange fluxes.

78 **Methods**

79 **2.1 Site description**

80 River Erpe is an urban lowland river, located at the eastern edge of Berlin, Germany, that receives
81 between 60% and 80% of its discharge from the large municipal WWTP Münchehofe. A detailed
82 description of river morphology and general biogeochemical traits of River Erpe can be found elsewhere
83 (Lewandowski et al., 2011). The study site is located at Heidemühle (Lat: 52.478647, Long: 13.635146)
84 roughly 1 km downstream of the confluence of River Erpe and the WWTP Münchehofe effluent.

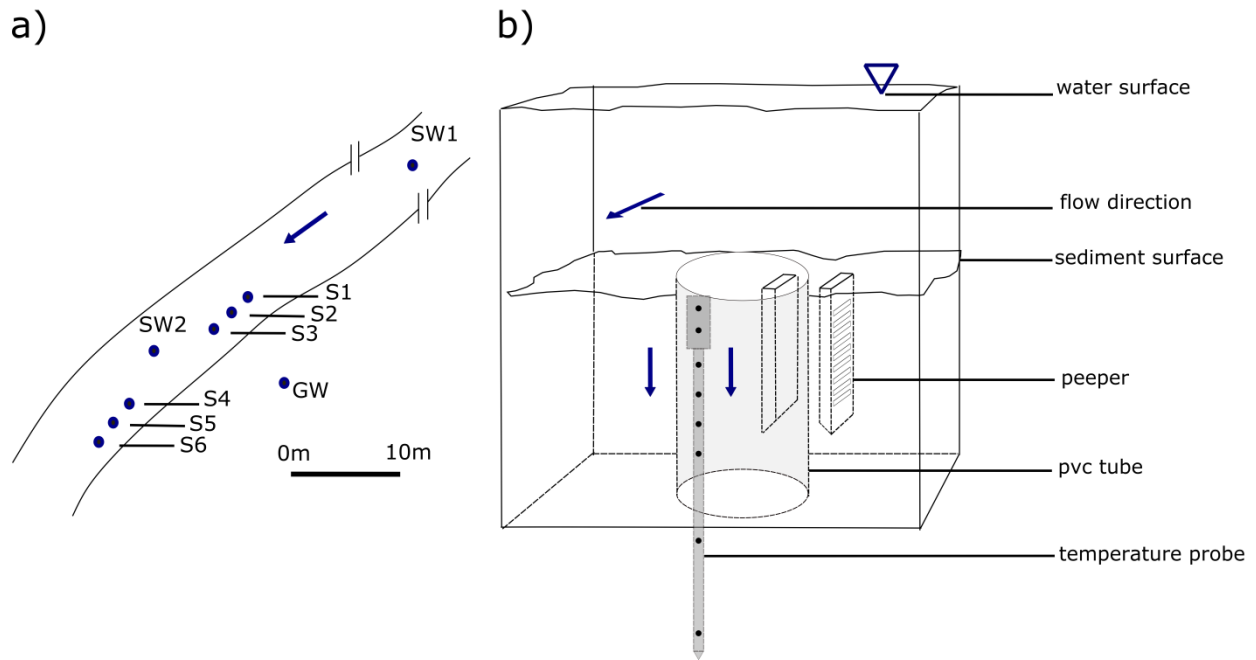
85 Streambed morphology at the field site was dominated by small dunes (crest height < 5 cm) and
86 macrophyte cover during the study period was < 10%. Streambed sediment characteristics were
87 determined from one 30 cm long core per station taken in close proximity (< 10 cm) to the sampling
88 station using a hand auger (9.0 cm ID). Saturated hydraulic conductivities (K) at 10 °C, measured using a
89 KSAT device (UMS Munich), were found to be log-normally distributed with a mean value of 5.0×10^{-5}
90 m s^{-1} and a standard deviation of $7.4 \times 10^{-5} \text{ m s}^{-1}$. Mean sediment porosity calculated from oven dried
91 samples was 0.39 ± 0.1 . Organic carbon content (f_{oc}), determined by loss on ignition, ranged between 0.5
92 and 6 wt.% and was found to be log-normally distributed with a mean value of $2.4 \pm 2.7 \text{ wt.}\%$.

93 **2.2 Experimental setup overview and field investigation periods**

94 The experimental setup consisted of six stations (S1 – S6) located close to the south-eastern bank of River
95 Erpe (Figure 1a). Each sampling station was equipped with a ring enclosure (PVC pipes with a diameter
96 of 15 cm and a height of 32 cm) which was inserted vertically into the sediment, with the top rim of the
97 pipe being level with the sediment surface. Concentrations of TrOCs in effluent-receiving streams are
98 typically variable in time as both concentrations of TrOCs in WWTP effluent itself as well as the relative
99 proportion of WWTP effluent to overall stream discharge may change on a sub-daily and daily basis
100 (Lewandowski et al., 2011). To overcome this challenge, we used dialysis chambers, also called peepers
101 (Hesslein, 1976), a passive sampling technique that integrates porewater solute concentrations over a time

102 span of several days and thus allows the sampling of time integrated, quasi-steady-state concentration
103 depth profiles in the hyporheic zone. A detailed description on the functionality of peepers and their
104 applicability to sample TrOCs in the hyporheic zone is provided in the Supplementary Material (SM).
105 Between December 1st and December 15th, 2015, peepers were used to collect porewater depth profiles in
106 the hyporheic zone (details below, section 2.4). To sample porewater both within the ring enclosure and
107 outside of the ring enclosure, one peeper was inserted into the sediment within the ring enclosure and one
108 inserted directly adjacent to the first one but outside of the ring enclosure (Figure 1 b). Between December
109 1st 2015 and April 15th, 2016, temperature probes (details below, section 2.3) were installed inside and
110 outside the ring enclosures. Vertical seepage flux inside the ring enclosure at S1, S2, S3, S4 and S6 as well
111 as outside ring enclosures at S3 was calculated from field temperature data measured between December
112 1st 2015 and December 15th 2015. Due to a limited availability of temperature probes, vertical seepage
113 fluxes inside ring enclosure at S5 and outside ring enclosures at S1, S2, S4, S5 and S6 were calculated
114 from temperature data collected between January 5th and April 15th 2016. This approach was justified as
115 (i) head gradients between river stage and the riparian aquifer were close to time invariant on timescales of
116 weeks and months and (ii) flux measurements made at the same station and peeper position between
117 December 2015 and April 2016 showed only little variation. A detailed summary of deployment periods
118 and the respective hydrodynamic conditions of the River Erpe between December 2015 and April 2016
119 can be found in the SM. The original purpose of the ring enclosures was to limit the influence of potential
120 horizontal subsurface flow in the hyporheic zone. However, flow investigations (sections 2.3.2 and 3.1)
121 did not indicate the presence of strong horizontal flow components outside the ring enclosures. As a result,
122 the same reactive transport model was used to evaluate concentration profiles both inside and outside the
123 ring enclosures. Stream stage, temperature and electrical conductivity (EC) were measured continuously
124 every 5 min between December 1st, 2015 and April 15th, 2016 using a CTD Diver (Van Essen Instruments
125 B.V., Netherlands) installed in a stilling well (SW1, Figure 1a). Between February 8th and April 15th,
126 additional head measurements were performed using a Pressure Diver (Van Essen Instruments B.V.,

127 Netherlands) in a second stilling well (SW2, Figure 1a) and a groundwater observation well (GW, Figure
128 1a). Head measurements were performed to gain general information on groundwater-surface water
129 interactions at the field site and qualitatively confirm the temperature based flux calculations (section 2.3).
130 Stream stage at SW2 during peeper deployment was inferred from correlations of stream stages at SW1
131 and SW2 between February 8th and April 15th 2016 (Figure SM-02).
132 Daily variations of stream stage and EC showed a periodic saw-toothed pattern with stage and EC troughs
133 in the early morning and peaks around noon and in the afternoon (Figure SM-08). Water levels in the river
134 were consistently above groundwater level, even during early morning stage troughs. As a consequence,
135 the groundwater-surface water interactions were characterized by downwelling conditions, a finding
136 which is in qualitative agreement with the flux calculations in section 2.3.



138

139 **Figure 1** a) Study site with stations (S1 – S6) close to south-eastern bank of River Erpe. Surface water stilling well 1 (SW1) is
 140 approximately 100 m upstream of S1 while surface water stilling well 2 (SW2) is located between S3 and S4. A groundwater well
 141 (GW) with filter screen 1.5 m below the water table is located 5 m beside the river. b) Each of the six stations S1 – S6 consists of
 142 a 32 cm long PVC pipe (ring enclosure) with its upper end being level with the sediment – water interface. Porewater samplers,
 143 so-called peepers, are placed next to each other inside and outside of the pipe, with the closed side of each peeper facing towards
 144 the PVC wall. Blue arrows indicate predominant flow directions.

145 2.3 Calculation of vertical seepage fluxes and transport parameters

146 2.3.1 Measurements of temperature time series

147 Temperature depth profiles were obtained using Multi Level Temperature Sticks (MLTS, UIT, Dresden),
 148 which are described in detail by Munz et al., (2011). In brief, MLTS sticks are polyoxymethylene sticks of
 149 a total length of 0.66 m and a diameter of 0.02 m, which are equipped with eight TSIC-506 temperature
 150 sensors (accuracy 0.07 °C between 5 °C and 45 °C). The probes were placed vertically in the streambed,
 151 with the head of the temperature stick being level with the sediment surface. Temperature time series
 152 were obtained for sediment depths 7, 17, 22, 27, 32, 42 and 62 cm.

153 2.3.2 Calculation of vertical seepage fluxes using VFLUX 2.0

154 Vertical seepage flux estimates based on continuous time series of temperature depth profiles were
155 evaluated using the MATLAB code VFLUX 2.0 (Gordon et al., 2012). In brief, VFLUX 2.0 utilizes
156 dynamic harmonic regression (DHR) (Young et al., 1999) to extract amplitude damping ratios (A_r) and
157 phase shifts ($\Delta\phi$) from two sinusoidal temperature time series measured at two different depths within the
158 stream sediment. DHR is a general case of the harmonic regression model in which the trigonometric
159 parameters describing the frequency components of the raw temperature time series may vary over time.
160 In the present study we used the combined $A_r\Delta\phi$ method introduced by McCallum et al., (2012) to
161 calculate seepage fluxes.

162 Vertical seepage flux was evaluated at 2 h intervals between sensors in 7 and 17, 17 and 27 cm depth
163 resulting in center-of-pair depths of 12 and 22 cm, thus covering the depth range of porewater sampling.
164 Monte Carlo Analysis was performed using the “vfluxmc” function implemented in VFLUX 2.0. Mean
165 seepage flux rates were calculated by first averaging flux over depth and then averaging depth-integrated
166 fluxes over time. To minimize errors induced by DHR filtering, the first and last 24 h of flux results were
167 discarded. Since VFLUX provides Darcy fluxes (q), calculated flux rates were divided by the average
168 porosity to derive porewater velocities (v).

169 2.3.3 Calculation of dispersion coefficients

170 Based on the calculated porewater velocities, the respective effective hydrodynamic dispersion
171 coefficients (D) were calculated via:

$$D = \tau D_{mol} + \alpha_L v \quad (1)$$

172
173
174 in which τ denotes the tortuosity, D_{mol} the molecular diffusion coefficient, v the porewater velocity and α_L
175 the longitudinal dispersivity. Tortuosity τ was assumed to be 0.7 and α_L was set to a tenth of the profile
176 length (2 cm). As D_{mol} for the respective TrOCs are generally smaller than $10^{-9} \text{ m}^2 \text{ s}^{-1}$ (Table SM-09) and

177 porewater velocities measured in the present study were larger than 10^{-6} m s⁻¹ (Table 1), the contribution
178 of D_{mol} to overall hydrodynamic dispersion coefficients was found to be negligible and thus molecular
179 diffusion was omitted in further calculations.

180 **2.4 Water sampling and analysis**

181 Deployed peepers (2.5 cm x 9.5 cm x 26.5 h x w x l) contained 20 chambers (approx. 10 mL of chamber
182 volume) in 1 cm increments. The chambers were filled with deionized water and covered with a
183 permeable membrane (polysulfon, 0.2 μm HT Tuffryn 200, Pall Corporation). Prior to installation, peepers
184 were placed into oxygen free water to reduce oxygen input into streambed sediments. In contrast to
185 common sampling practice the peepers had to be inserted completely into the sediment, because first trials
186 revealed a scouring around peepers if they were reaching into the overlying water. Thus, the peepers did
187 not sample overlying water or the immediate sediment-water interface. Peeper chambers were sampled on
188 site chamber by chamber immediately after recovery from the sediment using a needle and a 10 ml
189 syringe. Two ml of sampled porewater from each chamber were used for TrOC analysis, 2 ml were used
190 for nitrate analysis and 3 ml were used for major cation analysis. At SW1 an automatic water sampler
191 (model 3700, Teflon suction line, Teledyne ISCO, Lincoln, NE.) was deployed to take surface water
192 samples from December 1st to December 15th. Approximately 350 ml of surface water were collected
193 every hour from 20 cm below the water surface. Samples were combined in a flow proportional manner
194 based on stream stage derived hourly discharge, yielding a flow-weighted surface water sample for each
195 day during the peeper deployment period. Details on discharge measurements and the rating curve can be
196 found in the SM. Integrated water samples were used to measure time series of surface water pH, TrOCs
197 and major cation concentrations. Details on sample filtering and storage are provided in the SM.

198 The TrOCs studied include the pharmaceuticals gabapentin (GAB), bezafibrate (BZF), diclofenac (DCF),
199 venlafaxine (VLX), 4-formylaminoantipyrin (FAA), primidone (PRI), carbamazepine (CBZ), and
200 metoprolol (MTP) as well as the artificial sweetener acesulfame (ACS) and two benzotriazoles (1H-
201 Benzotriazole, BTA, and 5-Methylbenzotriazole, MBT), which are mostly used as anti-corrosion agents.

202 TrOCs were measured via high performance liquid chromatography with tandem mass spectrometry
203 (HPLC-MS/MS) following a method established previously (Zietzschmann et al., 2016, 2015). Separation
204 was achieved on a XSelect HSS T3 HPLC column (2.5 μm particle size, 2.1 * 50 mm, Waters, USA)
205 using a linear gradient (ultra pure water with 5 vol.-% methanol (HPLC grade, J.T. Baker, USA) and 0.1
206 vol.-% formic acid (HPLC grade, Sigma Aldrich, Germany) versus 100% methanol). A TSQ Vantage
207 mass spectrometer (Thermo Fisher Scientific, USA) in ESI \pm modes was used for compound analysis.
208 Each TrOC was identified by the characteristic ratio of the peak areas of two fragments, which were
209 chosen according to the DAIOS database (Zweckverband Landeswasserversorgung, Stuttgart, Germany).
210 For TrOC quantification, deuterated internal standards (Toronto Research Chemicals, Canada) were used
211 in combination with the more abundant fragment ion. Calibration curves consisted of 9 standards that
212 ranged from 0.01 to 55 $\mu\text{g/L}$ and were fitted using a weighted (1/x) linear regression model (coefficients
213 of determination, R^2 , were always > 0.99). Limits of quantification (LOQs, Table SM-01) were set to the
214 calibration level at which the less abundant fragment ion showed a signal to noise ratio of ≥ 3 . Data
215 evaluation was performed with Xcalibur 2.1 (Thermo Scientific). Additional details on the method are
216 provided in the SM.

217 Major cations and dissolved iron concentrations were determined via inductively coupled plasma optical
218 emission spectrometry (ICP-OES, ICP iCAP 6000series, Thermo Fisher). Nitrate-N was measured using a
219 continuous-flow-analyser (SAN ++, Skalar) after DIN EN ISO 13395. For each surface water sample, pH
220 was determined using a handheld pH meter, calibrated prior to each use (SenTix 41, WTW).

221 **2.5 Modeling of removal rate constants**

222 The one-dimensional advection-dispersion transport equation including first-order decay was used to
223 calculate removal rate constants of TrOCs from porewater depth profiles. Assuming steady-state
224 conditions in a homogeneous medium, negligible retardation of TrOCs along flow paths and uniform flow
225 conditions, the 1D advective-dispersion transport equation can be written as:

$$226 \quad 0 = D \frac{d^2c}{dx^2} - v \frac{dc}{dx} - kc \quad (2)$$

227
228 In which v denotes the vertical porewater velocity (m s^{-1}), D the effective hydrodynamic dispersion
229 coefficient ($\text{m}^2 \text{s}^{-1}$) and k the first-order removal rate constant (s^{-1}). Assuming that the solute concentration
230 is zero at infinite depth, the solution to Equation 4 can be written as:

$$231 \quad c(x) = c_{in} e^{\frac{v - \sqrt{v^2 + 4Dk}}{2D} x} \quad (3)$$

232
233 In which c_{in} denotes an arbitrary input concentration at the sediment-water interface. Following
234 linearization, the exponent of Equation 3 was found by fitting a linear model to the concentration data
235 using the statistical software package R (R Core Team, 2016). Details on the derivation of the analytical
236 solution (Equation 3) and parameterization can be found in the SM. It should be noted that the removal
237 rate constant (k), although mathematically treated as a first-order constant, actually represents a pseudo
238 constant, as it integrates a variety of attenuation processes. Removal rate constants were set to zero if the
239 relative removal within one profile or sub-profile was smaller than two times the associated standard
240 deviation. For calculation purposes, if the concentration of a compound in a depth profile was below the
241 LOQ, its concentration was set to the respective LOQ. Model assumptions are considered to be reasonably
242 well met as hydrological conditions and concentrations of TrOCs in the surface water remained relatively

243 constant over the peeper sampling period (Figures SM-08 & 11). Under these conditions peepers are likely
244 to sample quasi-steady-state concentration profiles in the HZ. As the majority of TrOCs investigated in the
245 present study are either neutral or negatively charged under ambient pH (pH = 7.7), electrostatic
246 interactions between TrOCs and hyporheic material are considered to be negligible. The investigated
247 TrOCs were further characterized by low $\log D_{ow}$ values (Table SM-10) and thus retardation of TrOCs in
248 the HZ due to partitioning into sediment organic matter is likely to be small. Even for compounds that
249 show comparatively high $\log D_{ow}$ values, such as CBZ and BTA, previous field studies found very low
250 retardation factors (Huntscha et al., 2013). Concentration time series of TrOCs in the surface water as well
251 as a detailed discussion of the relevance of sorption and the assumption of quasi-steady state are provided
252 in the SM. Since the study site was characterized by consistent downwelling (sections 2.2 and 3.1),
253 concentration data were not corrected for dilution with riparian groundwater. In order to compare removal
254 rate constants to literature values, compound half-lives were calculated according to:

$$t_{1/2} = \frac{\ln(2)}{k} \quad (4)$$

255

256 **3. Results and Discussion**

257 **3.1 Transport characteristics in the hyporheic zone**

258 The mean vertical porewater velocity (positive = downward) ranged from 0.12 m d⁻¹ inside S4 to 0.89 m d⁻¹
259 inside S2 (Table 1). Fluxes at S2, S3, S4, S5 and S6 showed no difference with peeper position (i.e.
260 inside and outside ring enclosures), while the flux calculated outside S1 was only half the flux calculated
261 inside S1.

262 **Table 1** Mean values ± standard deviations of porewater velocities (v) for the different sampling stations and inside/outside ring
263 enclosures. Porewater velocities were calculated by dividing VFLUX derived Darcy velocities by sediment porosity.

Station	v	
	m d ⁻¹	
	inside	outside
S1	0.84 ± 0.32	0.36 ± 0.13
S2	0.89 ± 0.31	0.83 ± 0.34
S3	0.84 ± 0.29	0.85 ± 0.30
S4	0.12 ± 0.05	0.25 ± 0.08
S5	0.59 ± 0.22	0.65 ± 0.24
S6	0.78 ± 0.29	0.81 ± 0.30

264

265 The mean vertical seepage fluxes measured inside and outside the ring enclosures were well within the
266 range reported in similar studies (Fitzgerald et al., 2015; Gordon et al., 2012) and matched downward
267 seepage rates previously calculated for the same river stretch (Lewandowski et al., 2011). During the
268 entire investigation period (December 2015 – April 2016) flux estimates indicated positive (downward)
269 flux which is consistent with the general head gradient at the study site, indicating downward flux from
270 the river to the riparian aquifer (section 2.2).

271 Errors associated with the calculation of vertical seepage fluxes using the McCallum $A_r\Delta\phi$ method mainly
272 result from non-ideal field conditions that violate model assumptions, such as non-uniform, non-steady
273 and multidimensional flow. As the ring enclosure only allows purely vertical and uniform flow, the model
274 assumptions of uniform one-dimensional flow fields are highly likely to be met for fluxes measured inside
275 ring enclosures. However, the flow field outside the ring enclosures could have been subjected to both
276 vertical and horizontal flow components. Additional model calculations indicated that horizontal
277 hyporheic exchange induced by small sand dunes and ripples was $< 0.05 \text{ m d}^{-1}$ (see SM for details).
278 Moreover, fluxes calculated inside ring enclosures resembled fluxes outside ring enclosures for all stations
279 but S1. It is thus unlikely that horizontal flow components significantly distorted the flow field outside
280 ring enclosures. Additional errors induced by non-steady flow conditions as well as non-ideal sensor
281 spacing and sensor accuracy were found to be of minor importance (see SM for detailed discussion).

282 We conclude that the vertical seepage fluxes calculated by the McCallum $A_r\Delta\phi$ method are likely to
283 resemble the actual vertical seepage fluxes quite well. Consequently, differences between sites and peeper
284 positions result likely from sediment heterogeneity and the associated variability in hydraulic
285 conductivity.

286 **3.2 Redox zonation in the HZ**

287 Nitrate concentrations were high in the surface water (median NO_3^- -N concentration 3.5 mg/L, Table SM-
288 07) and decreased rapidly within the streambed sediments. The only exception is the profile outside S2, in
289 which nitrate concentrations remained relatively constant. As nitrate was depleted, concentrations of total
290 dissolved iron and dissolved manganese (Fe_{tot} , Mn) increased (Figure SM-12). The depth ranges between
291 the sediment-water boundary and the depth at which nitrate was completely depleted as well as the range
292 between the depth at which Mn and Fe_{tot} concentrations increased and the lower end of the peeper were
293 considered to mark the extent of oxic to suboxic and anoxic redox conditions within the HZ (Figure SM-
294 12, Figure 2). The mean extent of oxic/suboxic and anoxic sub-sections in the HZ was 11 ± 4 cm and $10 \pm$
295 3 cm, respectively. The rapid consumption of surface water borne electron acceptors along relatively short
296 flow paths and hence the development of steep redox gradients at the surface water-streambed interface is
297 typical for hyporheic zones (Krause et al., 2017) and in agreement with previous studies on sediment
298 biogeochemistry conducted at River Erpe (Lewandowski et al., 2011).

299 The spatial extent of the redox zonation, i.e. the depth and steepness of the redoxcline, is generally a
300 function of both transport characteristics as well as biogeochemical parameters, such as the quality and
301 abundance of electron donors and acceptors, microbial turnover rates or the content of sediment organic
302 matter. In previous field investigations higher exchange fluxes led to deeper oxygen penetration depths
303 and later onset of denitrification (Harvey et al., 2013). The relatively steep redox gradients inside the ring
304 at S4 and S5 correspond qualitatively to lower seepage fluxes measured inside the ring enclosures at these
305 stations. However, redox profiles inside the ring enclosures at S1 and S2 as well as redox profiles inside
306 and outside S2, S5 and S6 were remarkably different, although vertical seepage rates were similar.
307 Overall, no quantitative correlation was found between calculated vertical seepage flux rates and the
308 extent of redox zones as shown in Figure 2. Since high resolution measurements of sedimentological
309 characteristics of Erpe sediment were not available, it is difficult to differentiate between the combined
310 effects of sedimentary and hydrological controls on redox zonation. The spatial variability of sediment

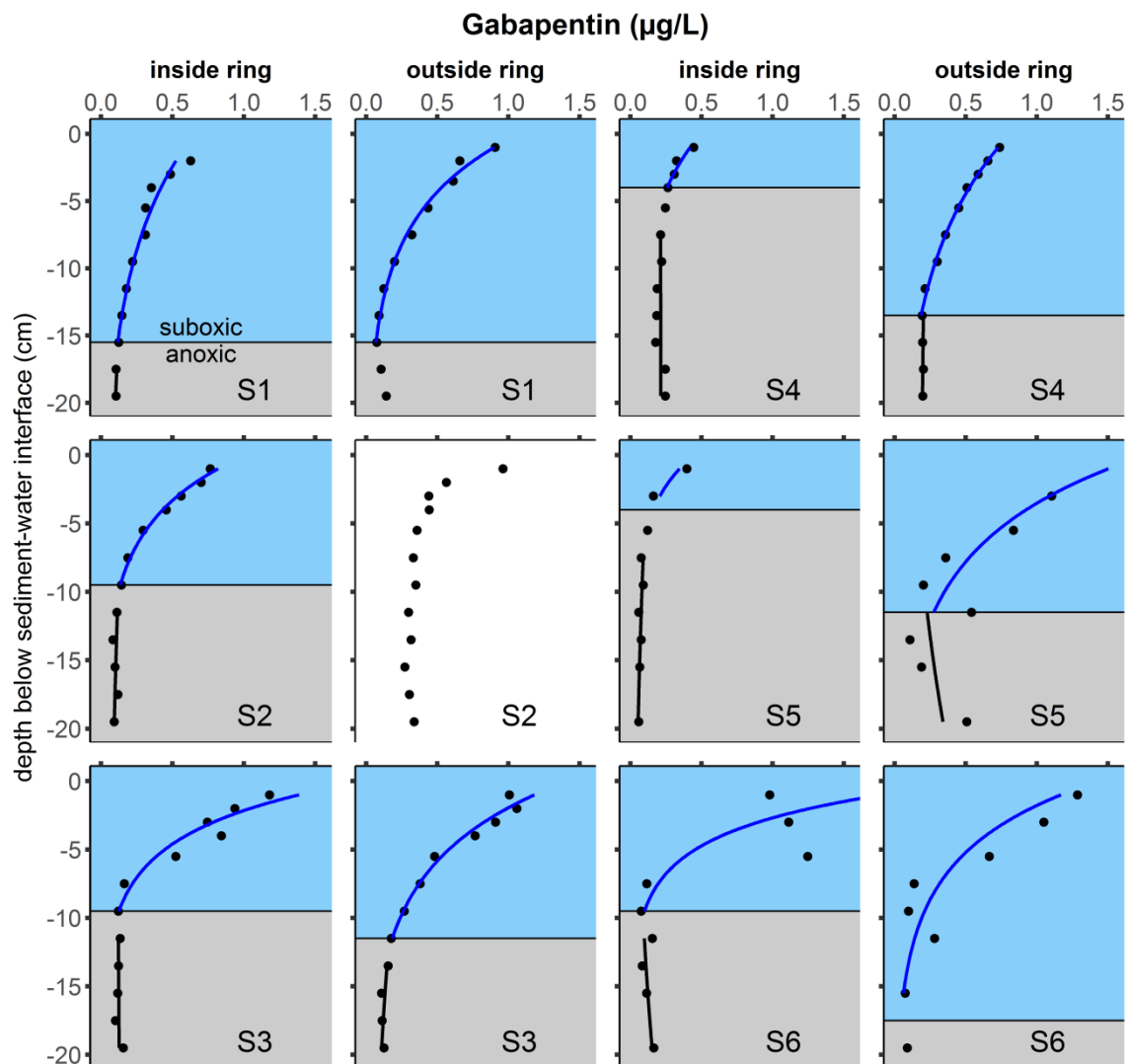
311 organic matter (f_{oc}) in the HZ of River Erpe, however, was relatively high (Figure SM-14). It therefore
312 seems likely that the spatial extent of redox zonation is not only a function of the magnitude of vertical
313 seepage flux but is also to a large extent controlled by the spatial distribution of sediment organic matter.

314 **3.3 Reactivity of TrOCs in the HZ**

315 Removal rate constants were calculated over the entire profile range and for each redox zone (suboxic and
316 anoxic) shown in Figure 2. Note that the profile taken outside S2 was excluded as oxic conditions
317 prevailed and no denitrification was observed. Kruskal-Wallis tests demonstrated that for all compounds
318 but SMX, DTA, CBZ and VLX removal rate constants derived for suboxic conditions were significantly
319 larger ($p < .05$) than the ones calculated for anoxic conditions (Table 2). Consequently, for TrOCs that
320 showed redox-dependent attenuation, removal rate constants are only reported for the respective redox
321 conditions. Median half-lives obtained from oxic/suboxic conditions varied between 0.43 d for PRI and
322 0.04 d for GAB (Table 2). General stability of TrOCs in the suboxic zone was in the order $GAB < ACS \approx$
323 $BZF < BTA < DCF < MBT < PRI$. Median half-lives for the same set of compounds calculated from
324 anoxic sections of the profile did not differ from zero or were even negative. Depth profiles of gabapentin,
325 associated model fits as well as the spatial extent of oxic/suboxic and anoxic zones are shown in Figure 2.
326 Depth profiles for the remaining TrOCs, individual removal rate constants calculated from every (sub-)
327 profile and a comparison of TrOC concentrations measured in the present study to concentrations
328 previously measured in River Erpe are provided in the SM. Due to analytical difficulties, removal rate
329 constants and respective half-lives for ACS and VLX were only calculated from eight profiles.

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332

333 **Figure 2** Equilibrium depth profiles of gabapentin (black dots) at the different sampling stations and inside/outside ring
 334 enclosures. The extent of oxic/suboxic and anoxic zones, as indicated by the presence of nitrate as well as dissolved Mn and Fe_{tot} ,
 335 within the profiles is indicated with blue and grey shading, respectively. Modelled concentration profiles are shown for both
 336 oxic/suboxic and anoxic conditions (solid lines). Note that for outside of the ring enclosure at S2 no removal rate constants were
 337 calculated as redox condition remained oxic throughout the profile and no denitrification occurred.

338 **Table 2** Average removal percentages calculated over the entire profile, median half-lives and associated interquartile ranges (IQR) for all TrOCs investigated in this study. Kruskal-
 339 Wallis test statistics indicate whether removal rate constants obtained for oxic/suboxic conditions differed significantly from removal rate constants obtained under anoxic conditions. *N*
 340 denotes the total number of profiles and *n* the number of profiles in which removal was insignificant. According to relative removal (%) and the redox dependency of removal rate
 341 constants, TrOCs were grouped into four groups: (i) removal > 15%, significant redox influence on removal rate constants (ii) removal > 15% , insignificant redox influence on
 342 removal rate constants (iii) removal ≤ 15%, insignificant redox influence on removal rate constants and group (iv) compounds that showed ambiguous behaviour. Note that individual
 343 removal rate constants calculated from every (sub-) profile as well as median surface water concentrations of the investigated TrOCs are provided in the SM.

TrOC	Relative removal (%)	Kruskal-Wallis test statistics		Median $t_{1/2} \pm$ IQR $t_{1/2}$		<i>n/N</i>	
		p - value	Chi square	oxic/suboxic	anoxic	oxic/suboxic	anoxic
Gabapentin ⁽ⁱ⁾ (GAB)	77±4	< 0.01	13.6	0.04±0.05	0.00±0.00	0/11	7/9
Bezafibrate ⁽ⁱ⁾ (BZF)	61±7	< 0.01	12.5	0.05±0.06	0.00±0.16	0/11	5/9
Acesulfame ⁽ⁱ⁾ (ACS)	66±7	< 0.01	8.0	0.05±0.02	-4.12±0.44	1/7	3/7
Benzotriazole ⁽ⁱ⁾ (BTA)	51±7	< 0.01	14.8	0.09±0.13	0.00±0.00	0/11	7/9
Diclofenac ⁽ⁱ⁾ (DCF)	36±9	< 0.01	9.7	0.16±0.13	0.00±0.00	3/11	8/9
Methylbenzotriazole ⁽ⁱ⁾ (MBT)	36±9	< 0.01	13.3	0.18±0.26	0.00±0.00	1/11	9/9
Primidon ⁽ⁱ⁾ (PRI)	32±10	< 0.01	7.9	0.43±0.14	0.00±0.00	4/11	9/9
4-Formylaminoantipyrin ^(iv) (FAA)	9±17	< 0.01	6.5	0.00±0.15	0.00±2.19	6/11	6/9
Metoprolol ^(iv) (MTP)	-6±17	0.04	4.1	0.23±0.14	0.00±0.43	3/11	4/9
				entire profile		entire profile	
Sulfamethoxazole ⁽ⁱⁱⁱ⁾ (SMX)	46±9	0.08	3.1	0.29±0.18		1/12	
Diatrizoic acid ⁽ⁱⁱⁱ⁾ (DTA)	33±10	0.10	2.7	0.32±0.22		3/12	
Venlafaxine ⁽ⁱⁱⁱ⁾ (VLX)	14±12	0.17	1.9	0.00±0.74		5/8	
Carbamazepine ⁽ⁱⁱⁱ⁾ (CBZ)	15±12	0.19	1.7	0.00±0.49		8/12	

344 Stability of TrOCs in the HZ varied widely with some TrOCs being readily attenuated in the HZ (i.e.
345 removal > 15% over the entire profile) and others being rather persistent (i.e. removal \leq 15%). Similar to
346 the findings of previous field and laboratory studies (Burke et al., 2014; Wiese et al., 2011), removal rate
347 constants for the majority of TrOCs investigated showed a strong dependence on redox conditions. Based
348 on the redox dependency as well as relative removal within the HZ, TrOCs were grouped into (i)
349 compounds that were removed by more than 15% and showed redox-dependent removal rate constants (ii)
350 compounds that were removed by more than 15% and showed redox-independent removal rate constants,
351 (iii) compounds that were removed by less than 15% and (iv) compounds that showed ambiguous
352 behavior within the sampled profiles.

353 Compounds in group i comprised ACS, GAB, BZF, BTA, MBT, DCF and PRI. Data on the fate of the
354 anticonvulsant gabapentin in saturated sediments is limited but a laboratory column study showed that
355 oxic conditions promote GAB removal (Hellauer et al., 2017). Acesulfame, an artificial sweetener, was
356 originally suggested to act as an ideal wastewater tracer (e.g., Engelhardt et al., 2013). More recent
357 studies, however, demonstrated as well as the present study that the compound undergoes redox-
358 dependent attenuation, with higher removal observed under oxic conditions (Burke et al., 2014; Regnery
359 et al., 2015). Reported half-lives for ACS under oxic/suboxic redox conditions range from 0.21 d to 4.2 d,
360 exceeding the median half-life found in the HZ of River Erpe (0.05 d) by more than one order of
361 magnitude. The use of ACS as a tracer is therefore not recommended, especially if oxic to suboxic
362 conditions might occur along the flow path. Reported half-lives of bezafibrate, a lipid lowering agent,
363 range from 1.0 d in the transient storage zone of a natural river (Riml et al., 2013) to 1.9 d in a laboratory
364 flume study (Li et al., 2015) and are thus almost two orders of magnitude higher than the median half-life
365 calculated for oxic/suboxic conditions in the present study (0.05 d). The anti-corrosive agents
366 benzotriazole and its methyl analogue 5-methyl-1H-benzotriazole have both been shown to be relatively
367 stable in saturated sediments. (Liu et al., 2013), for instance, reported redox-dependent removal for MBT
368 and BTA in microcosms with half-lives of 60 d and 76 d for nitrate reducing conditions, while BTA was

369 found to be persistent in a column study that used hyporheic sediment (Burke et al., 2014). These findings
370 are in contrast to the data reported in the present study showing that both BTA and MBT undergo rapid
371 removal within the HZ. Among all investigated TrOCs, BTA and MBT were the only ones for which
372 concentrations strongly correlated with each other (adj. $R^2 = 0.90$, p -value < 0.01), indicating that, in
373 principle, the same process seems to affect the removal of both compounds. A general redox dependency
374 of the removal of diclofenac, an analgesic, has been shown in both laboratory and field studies (Burke et
375 al., 2014; Wiese et al., 2011). Under oxic/suboxic conditions, reported half-lives ranged from 3 d to 36 d
376 within bank filtration settings (Henzler et al., 2014; Wiese et al., 2011). Shorter half-lives (0.13 d) were
377 observed in a column study (Burke et al., 2014), a value that agrees well with the half-life found in the
378 present study for oxic/suboxic conditions. Although the concentrations of primidone, an anticonvulsant,
379 did not decrease significantly in some of the oxic/suboxic sub-profiles (4 out of 11 profiles), overall mean
380 removal was found to be $32 \pm 10\%$. This was not expected as previous studies found half-lives for PRI in
381 the order of several years (Regnery et al., 2015).

382 Compounds of group ii comprised the anti-bacterial agent sulfamethoxazole as well as the X-ray contrast
383 agent diatrizoic acid. Both compounds were attenuated in the HZ, but calculated removal rate constants
384 did not depend on redox conditions. For SMX, this finding is in agreement with previous findings
385 indicating its removal under both oxic and anoxic conditions (Baumgarten et al., 2011). In a laboratory
386 flume study and during bank filtration, half-lives of 34 d and 22 d were reported for SMX respectively
387 (Henzler et al., 2014; Li et al., 2015) exceeding the median half-life found in this study (0.29 d) by
388 roughly two orders of magnitude. Removal of DTA during wastewater treatment was reported (Haiß and
389 Kümmerer, 2006) but limited data on removal rate constants in natural sediments is available.

390 The anticonvulsant carbamazepine and the neuro-active compound venlafaxine (group iii) were found to
391 be rather stable in the HZ of River Erpe. CBZ was previously reported to be very persistent in saturated
392 sediments (Bertelkamp et al., 2014; Henzler et al., 2014). Although removal of CBZ has been described
393 for iron-reducing conditions (Wiese et al., 2011), ambient exposure times to such redox conditions in the

394 present study were likely too short for the compound to be significantly removed. For VLX removal has
395 been demonstrated in the laboratory column study by Hellauer et al., (2017), but the compound was found
396 to be persistent in streams and during river bank filtration (Huntscha et al., 2013; Writer et al., 2013). It is
397 likely that residence times in the investigated section of the HZ were too short to achieve significant
398 attenuation for CBZ and VLX. Although median half-lives were zero for both compounds, in some
399 profiles concentrations of VLX and CBZ decreased significantly allowing the calculation of
400 comparatively low removal rate constants.

401 Similar inconsistent behaviour was also found for the compounds of group iv, the beta-blocker metoprolol
402 and 4-formylaminoantipyrin, a metabolite of the antipyretic drug dipyrone. Although MTP was persistent
403 in the transient storage zone of a natural river (Riml et al., 2013), half-lives of several hours were reported
404 in laboratory studies (Bertelkamp et al., 2014; Burke et al., 2014). Removal of 4-formylaminoantipyrin in
405 saturated sediments has been demonstrated in both laboratory and field studies with half-lives ranging
406 from 0.12 to 14 days (Burke et al., 2014; Wiese et al., 2011). Although overall removal for FAA and MTP
407 was insignificant, both showed positive removal during oxic/suboxic conditions and negative removal
408 rates (indicating concentrations increase) during anoxic conditions.

409 The order of relative stability of TrOCs encountered in the HZ of River Erpe partly differs from relative
410 stabilities described in other studies. For instance, CBZ was often found to be relatively more stable than
411 DCF (Burke et al., 2014; Wiese et al., 2011), which is in accordance with the findings of the present study.
412 In contrast, ACS has been found to be more stable than DCF (Burke et al., 2014; Regnery et al., 2015),
413 although in the present study DCF was more stable than ACS. Unsystematic findings with respect to
414 relative stabilities of TrOCs were mainly attributed to sediment heterogeneity and microbial community
415 diversity (Radke and Maier, 2014).

416 It should be further noted, that the calculation of removal rate constants in the present study is based on
417 field data collected under winter conditions with average hyporheic temperatures of 10°C (Figure SM-13).
418 Previous studies have shown that with increasing temperatures, removal rate constants increase (Burke et

419 al., 2014). However, with increasing temperatures redox zonation may also shift upwards thus limiting
420 exposure time to oxic/suboxic redox conditions and rendering the HZ less efficient in removing
421 compounds that show preferred removal under oxic/suboxic conditions.

422 In summary, in-situ removal rate constants of the majority of TrOCs in the HZ of River Erpe are one or
423 even two orders of magnitude higher compared to removal rates reported in previous studies. We found
424 that even compounds that were previously regarded as very stable, such as BTA and PRI, are retained in
425 the HZ. As sorption is generally not regarded to considerably influence removal of polar TrOCs (see SM
426 for detailed discussion and $\log D_{ow}$ values), biotransformation is likely to be the dominant removal
427 mechanism for TrOCs in the HZ. Previous studies indicated that removal rate constants correlated
428 positively with microbial activity and TrOCs concentrations (Bertelkamp et al., 2016). The relatively high
429 removal rate constants in the HZ may thus be the results of site-specific characteristics of River Erpe, i.e.
430 high nutrient concentrations and high loads of TrOCs, high microbial activity in the upper centimetres of
431 the HZ as indicated by the steep redoxclines and the presence of a well-adapted microbial community.

432 **4. Conclusion**

433 The present study provides in-situ removal rate constants of several polar TrOCs in the HZ of an urban
434 lowland stream. Compared to laboratory studies and bank filtration settings, removal rate constants of
435 TrOCs in the HZ were relatively high and thus HZs may efficiently remove TrOCs along short flow paths.
436 River restoration measures that explicitly aim at increasing hyporheic exchange in streams that are
437 dominated by WWTP effluent may thus improve whole-stream TrOCs removal and hence water quality.
438 We demonstrate that removal efficiency of the HZ is controlled by both residence times (i.e. advective
439 transport velocity and the magnitude of groundwater-surface water exchange fluxes) and the reactivity of
440 the respective TrOCs. For the majority of TrOCs, removal rate constants and thus removal efficiency of
441 the HZ were significantly higher under oxic/suboxic conditions than under anoxic conditions. When
442 predicting the fate of TrOCs in the HZ or assessing the relevance of the HZ with respect to whole-stream
443 TrOCs attenuation, future studies should therefore consider exposure times to ambient redox conditions
444 rather than overall residence time distributions in the HZ. For TrOCs that show redox-dependent
445 biotransformation, relative removal may be increased if exposure times to oxic/suboxic conditions are
446 enlarged.

447 The present study also suggests that high microbial activity promotes TrOCs removal. However, high
448 microbial activity also leads to steep redox gradients and, for many TrOCs, is thus expected to limit the
449 spatial extent of reactive zones in the HZ. More research is needed to understand the in-situ interactions
450 between exchange fluxes, microbial metabolism and the extent of favourable “reactive” zones in the HZ.
451 While we provide insights into the general fate of parent TrOCs in the HZ, future studies should also
452 include transformation products and thus aim at closing mass balances along flow paths.

453 **References**

- 454 Baumgarten, B., Jählig, J., Reemtsma, T., Jekel, M., 2011. Long term laboratory column experiments to
 455 simulate bank filtration: Factors controlling removal of sulfamethoxazole. *Water Res.* 45, 211–
 456 220. <https://doi.org/10.1016/j.watres.2010.08.034>
- 457 Bertelkamp, C., Reungoat, J., Cornelissen, E.R., Singhal, N., Reynisson, J., Cabo, A.J., van der Hoek, J.P.,
 458 Verliefe, A.R.D., 2014. Sorption and biodegradation of organic micropollutants during river
 459 bank filtration: A laboratory column study. *Water Res.* 52, 231–241.
 460 <https://doi.org/10.1016/j.watres.2013.10.068>
- 461 Bertelkamp, C., van der Hoek, J.P., Schoutteten, K., Hulpiau, L., Vanhaecke, L., Vanden Bussche, J.,
 462 Cabo, A.J., Callewaert, C., Boon, N., Löwenberg, J., Singhal, N., Verliefe, A.R.D., 2016. The
 463 effect of feed water dissolved organic carbon concentration and composition on organic
 464 micropollutant removal and microbial diversity in soil columns simulating river bank filtration.
 465 *Chemosphere* 144, 932–939. <https://doi.org/10.1016/j.chemosphere.2015.09.017>
- 466 Boano, F., Harvey, J.W., Marion, A., Packman, A.I., Revelli, R., Ridolfi, L., Wörman, A., 2014.
 467 Hyporheic flow and transport processes: Mechanisms, models, and biogeochemical implications.
 468 *Rev. Geophys.* 52, 603–679. <https://doi.org/10.1002/2012RG000417>
- 469 Burke, V., Greskowiak, J., Asmuß, T., Bremermann, R., Taute, T., Massmann, G., 2014. Temperature
 470 dependent redox zonation and attenuation of wastewater-derived organic micropollutants in the
 471 hyporheic zone. *Sci. Total Environ.* 482, 53–61. <https://doi.org/10.1016/j.scitotenv.2014.02.098>
- 472 Engelhardt, I., Prommer, H., Moore, C., Schulz, M., Schüth, C., Ternes, T.A., 2013. Suitability of
 473 temperature, hydraulic heads, and acesulfame to quantify wastewater-related fluxes in the
 474 hyporheic and riparian zone. *Water Resour. Res.* 49, 426–440.
 475 <https://doi.org/10.1029/2012WR012604>
- 476 Fitzgerald, A., Roy, J.W., Smith, J.E., 2015. Calculating discharge of phosphorus and nitrogen with
 477 groundwater base flow to a small urban stream reach. *J. Hydrol.* 528, 138–151.
 478 <https://doi.org/10.1016/j.jhydrol.2015.06.038>
- 479 Gordon, R.P., Lautz, L.K., Briggs, M.A., McKenzie, J.M., 2012. Automated calculation of vertical pore-
 480 water flux from field temperature time series using the VFLUX method and computer program. *J.*
 481 *Hydrol.* 420, 142–158. <https://doi.org/10.1016/j.jhydrol.2011.11.053>
- 482 Greskowiak, J., Hamann, E., Burke, V., Massmann, G., 2017. The uncertainty of biodegradation rate
 483 constants of emerging organic compounds in soil and groundwater – A compilation of literature
 484 values for 82 substances. *Water Res.* 126, 122–133. <https://doi.org/10.1016/j.watres.2017.09.017>
- 485 Haiß, A., Kümmerer, K., 2006. Biodegradability of the X-ray contrast compound diatrizoic acid,
 486 identification of aerobic degradation products and effects against sewage sludge micro-organisms.
 487 *Chemosphere* 62, 294–302. <https://doi.org/10.1016/j.chemosphere.2005.05.007>
- 488 Harvey, J.W., Böhlke, J.K., Voytek, M.A., Scott, D., Tobias, C.R., 2013. Hyporheic zone denitrification:
 489 Controls on effective reaction depth and contribution to whole-stream mass balance: Scaling
 490 hyporheic flow controls on stream denitrification. *Water Resour. Res.* 49, 6298–6316.
 491 <https://doi.org/10.1002/wrcr.20492>
- 492 Hellauer, K., Mergel, D., Ruhl, A., Filter, J., Hübner, U., Jekel, M., Drewes, J., 2017. Advancing
 493 sequential managed aquifer recharge technology (SMART) using different intermediate oxidation
 494 processes. *Water* 9, 221. <https://doi.org/10.3390/w9030221>
- 495 Henzler, A.F., Greskowiak, J., Massmann, G., 2014. Modeling the fate of organic micropollutants during
 496 river bank filtration (Berlin, Germany). *J. Contam. Hydrol.* 156, 78–92.
 497 <https://doi.org/10.1016/j.jconhyd.2013.10.005>
- 498 Hesslein, R.H., 1976. An in situ sampler for close interval pore water studies. *Limnol. Oceanogr.* 21, 912–
 499 914.

500 Hoppe-Jones, C., Dickenson, E.R.V., Drewes, J.E., 2012. The role of microbial adaptation and
501 biodegradable dissolved organic carbon on the attenuation of trace organic chemicals during
502 groundwater recharge. *Sci. Total Environ.* 437, 137–144.
503 <https://doi.org/10.1016/j.scitotenv.2012.08.009>

504 Huntscha, S., Rodriguez Velosa, D.M., Schroth, M.H., Hollender, J., 2013. Degradation of polar organic
505 micropollutants during riverbank filtration: Complementary results from spatiotemporal sampling
506 and push–pull tests. *Environ. Sci. Technol.* 47, 11512–11521. <https://doi.org/10.1021/es401802z>

507 Krause, S., Lewandowski, J., Grimm, N.B., Hannah, D.M., Pinay, G., McDonald, K., Martí, E., Argerich,
508 A., Pfister, L., Klaus, J., Battin, T., Larned, S.T., Schelker, J., Fleckenstein, J., Schmidt, C.,
509 Rivett, M.O., Watts, G., Sabater, F., Sorolla, A., Turk, V., 2017. Ecohydrological interfaces as hot
510 spots of ecosystem processes. *Water Resour. Res.* 53, 6359–6376.
511 <https://doi.org/10.1002/2016WR019516>

512 Lewandowski, J., Putschew, A., Schwesig, D., Neumann, C., Radke, M., 2011. Fate of organic
513 micropollutants in the hyporheic zone of a eutrophic lowland stream: Results of a preliminary
514 field study. *Sci. Total Environ.* 409, 1824–1835. <https://doi.org/10.1016/j.scitotenv.2011.01.028>

515 Li, Z., Sobek, A., Radke, M., 2015. Flume experiments to investigate the environmental fate of
516 pharmaceuticals and their transformation products in streams. *Environ. Sci. Technol.* 49, 6009–
517 6017. <https://doi.org/10.1021/acs.est.5b00273>

518 Liu, Y.-S., Ying, G.-G., Shareef, A., Kookana, R.S., 2013. Biodegradation of three selected benzotriazoles
519 in aquifer materials under aerobic and anaerobic conditions. *J. Contam. Hydrol.* 151, 131–139.
520 <https://doi.org/10.1016/j.jconhyd.2013.05.006>

521 McCallum, A.M., Andersen, M.S., Rau, G.C., Acworth, R.I., 2012. A 1-D analytical method for
522 estimating surface water-groundwater interactions and effective thermal diffusivity using
523 temperature time series. *Water Resour. Res.* 48. <https://doi.org/10.1029/2012WR012007>

524 Munz, M., Oswald, S.E., Schmidt, C., 2011. Sand box experiments to evaluate the influence of subsurface
525 temperature probe design on temperature based water flux calculation. *Hydrol. Earth Syst. Sci.* 15,
526 3495–3510. <https://doi.org/10.5194/hess-15-3495-2011>

527 Pal, A., He, Y., Jekel, M., Reinhard, M., Gin, K.Y.-H., 2014. Emerging contaminants of public health
528 significance as water quality indicator compounds in the urban water cycle. *Environ. Int.* 71, 46–
529 62. <https://doi.org/10.1016/j.envint.2014.05.025>

530 R Core Team, 2016. R: A language and environment for statistical computing. R Foundation for Statistical
531 Computing, Vienna, Austria. URL <https://www.R-project.org/>.

532 Radke, M., Maier, M.P., 2014. Lessons learned from water/sediment-testing of pharmaceuticals. *Water*
533 *Res.* 55, 63–73. <https://doi.org/10.1016/j.watres.2014.02.012>

534 Regnery, J., Wing, A.D., Alidina, M., Drewes, J.E., 2015. Biotransformation of trace organic chemicals
535 during groundwater recharge: How useful are first-order rate constants? *J. Contam. Hydrol.* 179,
536 65–75. <https://doi.org/10.1016/j.jconhyd.2015.05.008>

537 Riml, J., Wörman, A., Kunkel, U., Radke, M., 2013. Evaluating the fate of six common pharmaceuticals
538 using a reactive transport model: Insights from a stream tracer test. *Sci. Total Environ.* 458–460,
539 344–354. <https://doi.org/10.1016/j.scitotenv.2013.03.077>

540 Schwarzenbach, R.P., Escher, B.I., Fenner, K., Hofstetter, T.B., Johnson, A.C., von Gunten, U., Wehrli,
541 B., 2006. The Challenge of Micropollutants in Aquatic Systems. *Science* 313, 1072–1077.
542 <https://doi.org/10.1126/science.1127291>

543 Wiese, B., Massmann, G., Jekel, M., Heberer, T., Dünnebier, U., Orlikowski, D., Grützmaier, G., 2011.
544 Removal kinetics of organic compounds and sum parameters under field conditions for managed
545 aquifer recharge. *Water Res.* 45, 4939–4950. <https://doi.org/10.1016/j.watres.2011.06.040>

546 Writer, J.H., Antweiler, R.C., Ferrer, I., Ryan, J.N., Thurman, E.M., 2013. In-Stream Attenuation of
547 Neuro-Active Pharmaceuticals and Their Metabolites. *Environ. Sci. Technol.* 47, 9781–9790.
548 <https://doi.org/10.1021/es402158t>

549 Young, P.C., Pedregal, D.J., Tych, W., 1999. Dynamic harmonic regression. *J. Forecast.* 18, 369–394.

550 Zietzschmann, F., Aschermann, G., Jekel, M., 2016. Comparing and modeling organic micro-pollutant
551 adsorption onto powdered activated carbon in different drinking waters and WWTP effluents.
552 Water Res. 102, 190–201. <https://doi.org/10.1016/j.watres.2016.06.041>
553 Zietzschmann, F., Mitchell, R.-L., Jekel, M., 2015. Impacts of ozonation on the competition between
554 organic micro-pollutants and effluent organic matter in powdered activated carbon adsorption.
555 Water Res. 84, 153–160. <https://doi.org/10.1016/j.watres.2015.07.031>
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